

Carbon Dioxide Transport in the Stable Boundary Layer over Heterogeneous Surfaces: A Large-Eddy Simulation Study

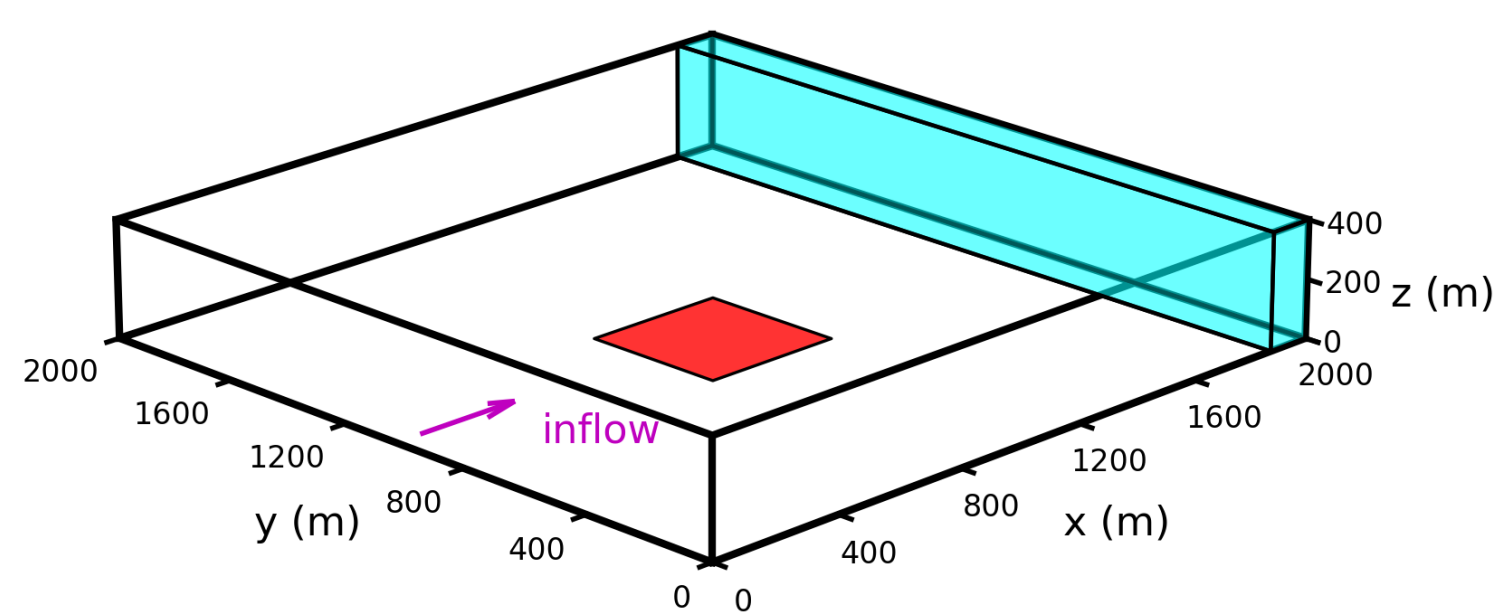
Shenghao Zhang, Johan Meyers

Funded by Agroflux Project

1. Introduction

- CO₂ fluxes in the nocturnal stable boundary layer are strongly influenced by surface heterogeneity
- Nine idealized LES cases are simulated to study the effects of thermal and flux heterogeneity
- Focus on dispersive motions and CO₂ budget closure

2. Numerical Aspect



- Stably stratified flow (≈ 2.2 K potential temperature difference)
- Heterogeneous surface patch (400×400 m²)
- Subject to atmospheric subsidence
- Concurrent precursor and fringe regions are applied

Simulated cases

$$T = \frac{T_{patch} - T_{ground}}{T_{top} - T_{ground}}$$

$$Q = \frac{Q_{patch}}{Q_{ground}}$$

Q / T	1	0.5	0.25
1	Q1T1	Q1T05	Q1T025
2	Q2T1	Q2T05	Q2T025
4	Q4T1	Q4T05	Q4T025

$$\frac{\partial \tilde{u}_i}{\partial x_i} = 0$$

$$\frac{\partial \tilde{u}_i}{\partial t} = -\frac{\partial}{\partial x_j} (\tilde{u}_i \tilde{u}_j + \tau_{ij}^{sgs}) - \frac{\partial p^*}{\partial x_i} + f_c \epsilon_{ij3} (\tilde{u}_j - G \delta_{j1}) + \frac{g}{\theta_r} (\tilde{\theta}_\delta - \langle \tilde{\theta}_\delta \rangle) \delta_{i3}$$

$$\frac{\partial \tilde{\theta}_\delta}{\partial t} = -\frac{\partial}{\partial x_j} (\tilde{u}_i \tilde{\theta}_\delta + q_j^{sgs}) + \Gamma_w x_3 \frac{\partial \langle \tilde{\theta}_\delta \rangle}{\partial x_3}$$

$$\frac{\partial \tilde{c}_\delta}{\partial t} = -\frac{\partial}{\partial x_j} (\tilde{u}_i \tilde{c}_\delta + q_{cj}^{sgs}) + \Gamma_w x_3 \frac{\partial \langle \tilde{c}_\delta \rangle}{\partial x_3}$$

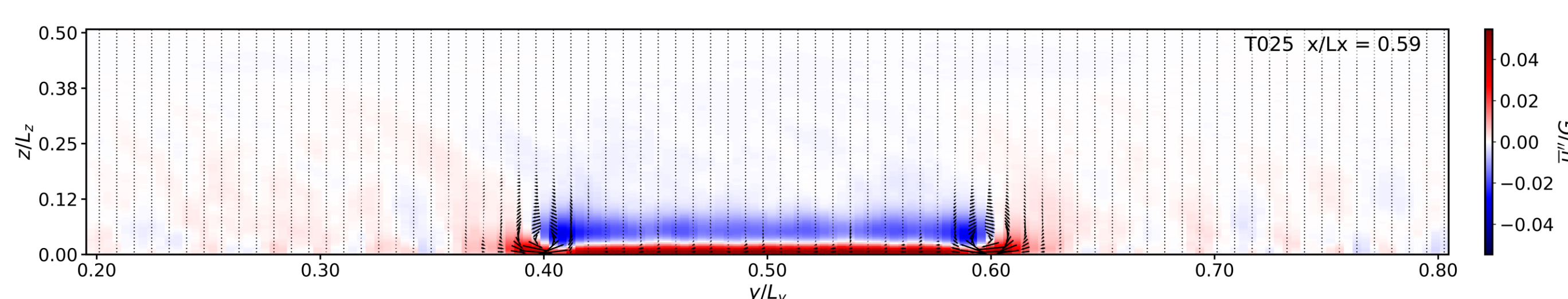
The Navier–Stokes and transport equations are solved using the SP-Wind solver

- In-house DNS/LES solver for atmospheric boundary layer flows
- Pseudo-spectral scheme in horizontal directions and 4th-order finite-difference scheme vertically
- Efficiently parallelized up to 10,000 CPUs on the high-performance infrastructure of the Flemish Supercomputer Center

3. Results

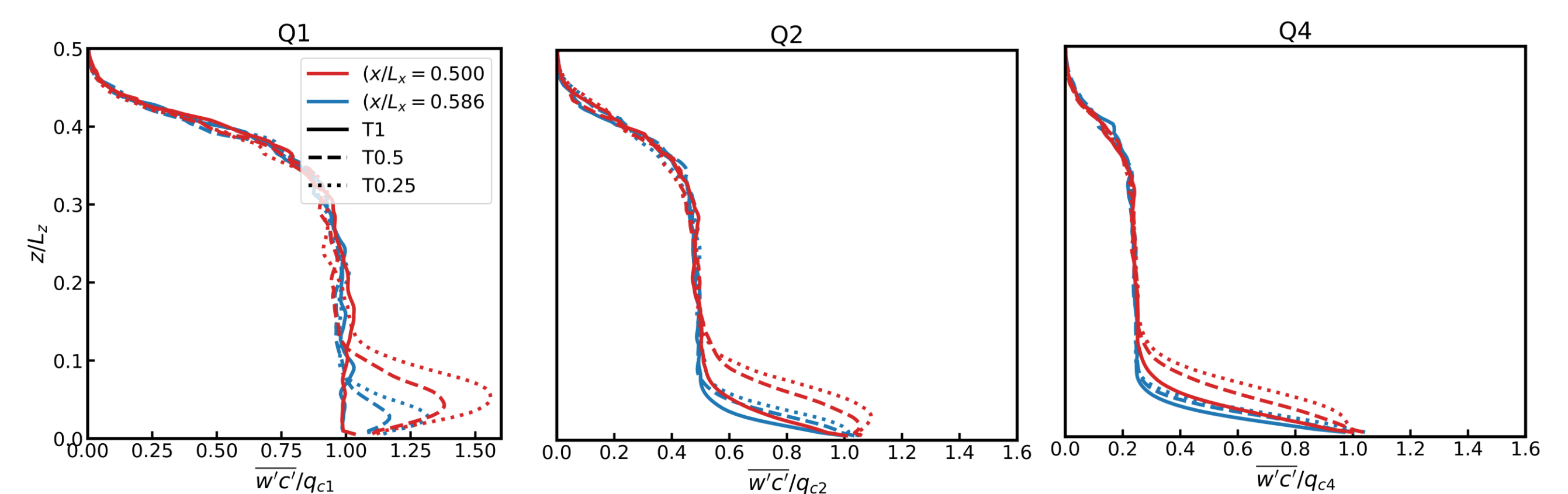
3.1 Dispersive velocities

$$\phi(x, y, z, t) = \langle \bar{\phi} \rangle(z) + \bar{\phi}''(x, y, z) + \phi'(x, y, z, t)$$



- Buoyancy effects at the edges generate swirls, resulting in LMP–HMP structures
- A second pair of LMPs–HMPs in vertical direction due to turbulent vertical momentum mixing

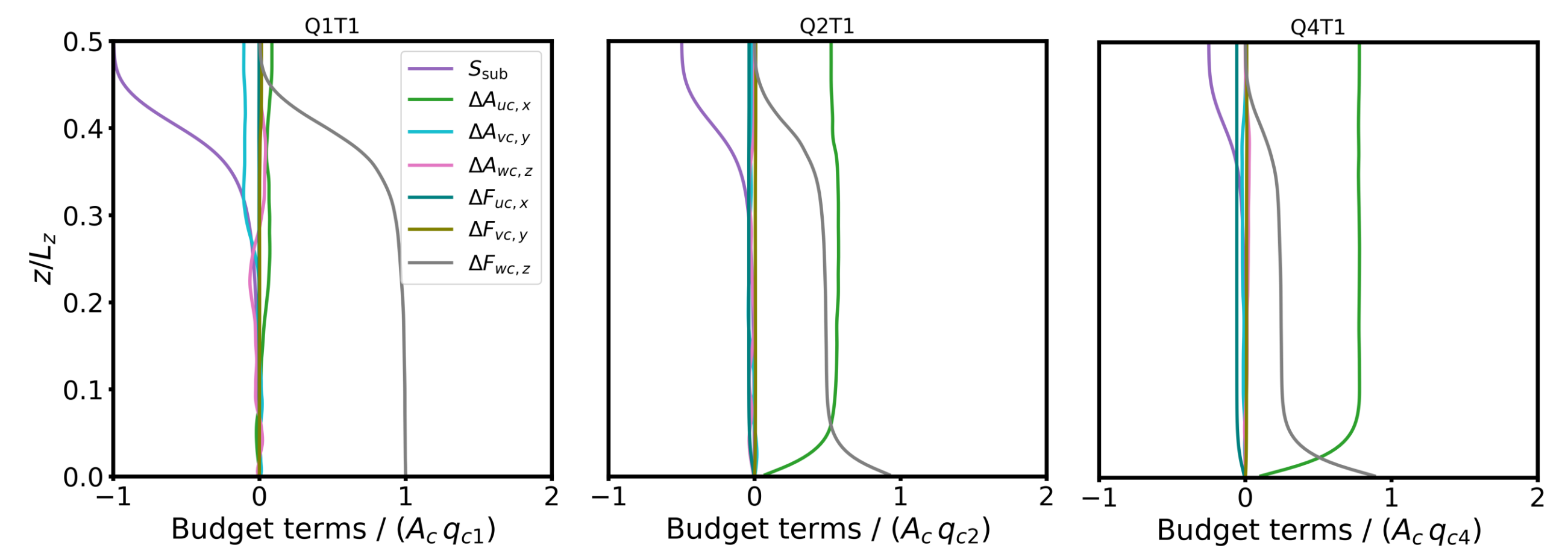
3.2 CO₂ turbulent vertical transport



- Both higher CO₂ flux and higher temperature increases vertical turbulent transport

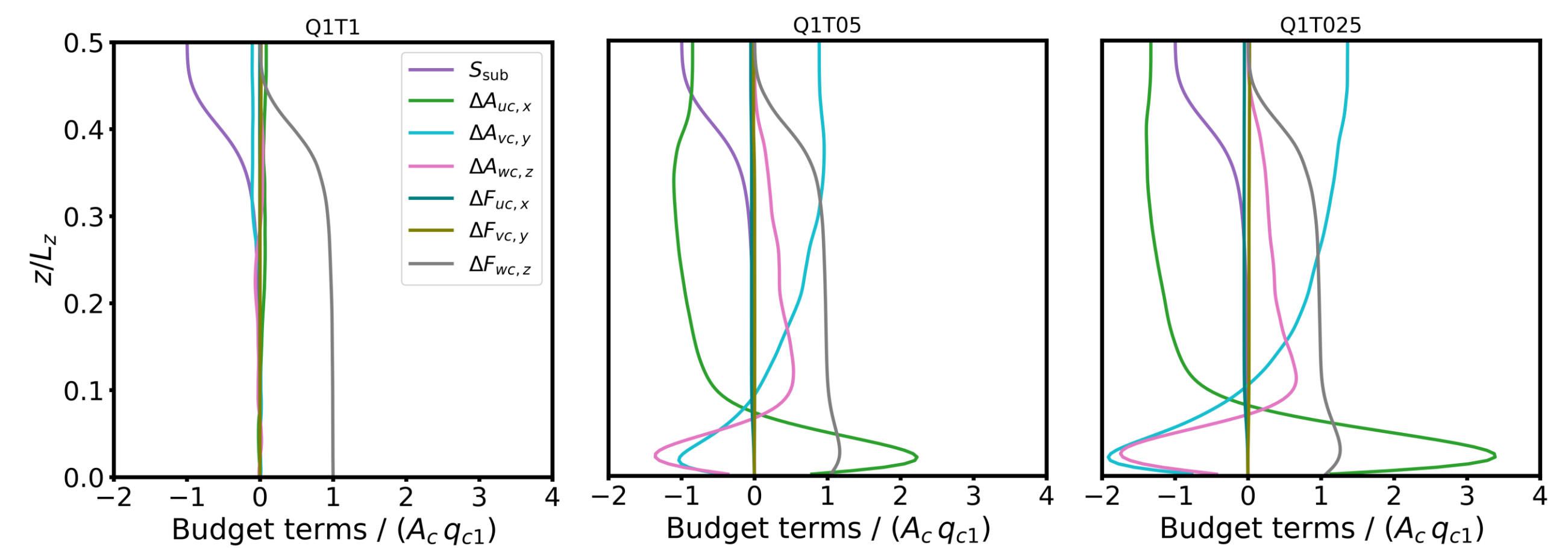
3.3 CO₂ Budget Analysis

3.3.1 Homogeneous T heterogeneous Q



- Under homogeneous condition, subsidence balances with vertical turbulent CO₂ transport
- When CO₂ flux is heterogeneous, balance is reached among subsidence, vertical turbulent transport and subsidence

3.3.2 Heterogeneous T homogeneous Q



- Heat-induced dispersive motions strongly alter local CO₂ balance components
- Clear edge effects when using control volumes right above the patch

4. Discussion

Surface heterogeneity reshapes CO₂ transport through coupled dispersive and turbulent processes, with their relative importance depending on thermal and flux contrasts.

However, Quantifying the relative contributions of dispersive and turbulent fluxes remains challenging. Case-by-case evaluation is needed under different atmospheric stability conditions.